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Influence of Meteorology on Surface Ozone, NO2 and Fine Particulate Matter at a Tropical Site in India

Abstract

Regular measurements of surface Ozone, NO2, PM2.5 and PM10 at Agra (27°10' N, 78°05' E, 169 msl), India have been carried out from Jan'11 to Dec'11using online O3 and NOx analyzers while PM2.5 and PM₁₀ were collected using high volume sampler. Simultaneously, temperature, solar radiation, wind speed and relative humidity were also recorded using automatic weather station (WM 251). The diurnal variation of O₃ showed noontime maxima and NO₂ displayed its maxima in morning and evening hours. The seasonal variation of O_3 and particulate matter (PM) showed summer and winter maxima respectively. The relationship between atmospheric pollutants and meteorological parameters was studied using correlation analysis. O₃ showed significant positive correlation with temperature and solar radiation. NO2 was found to be moderately correlated with wind speed while PM2.5 and PM10 were found to be significantly negatively correlated with temperature. Back trajectory analysis revealed that during summer and monsoon particulate matter at the site was contributed from distant sources while during the winter local sources appeared to dominate.

Keywords: Ozone, Oxides of Nitrogen, Particulate Matter, Meteorology, Back Trajectory

Introduction

The increasing concentrations of particulate matter (PM) and tropospheric ozone (O₃) in the ambient air are responsible for decreasing visibility in a region. Studies have established that exposure to ozone and particulate matter is harmful to human health (NRC 1998). Particulate matter (PM) occurs in the atmosphere due to emissions from a number of anthropogenic sources like gasoline and diesel powered vehicles, combustible fuels such as coal, wood, gas and oil, various industrial processes and volatilization from polluted ground. The main concern for particulate matter is also due to the presence of polycyclic aromatic hydrocarbons (PAH) possessing carcinogenic and mutagenic properties. O₃ is regarded as a powerful green house gas in the atmosphere after water vapor, CO₂, CH₄ and nitrous oxide with a radiative forcing of +0.35 Wm⁻² (IPCC 2001). O₃ is both cyto-toxic and phyto-toxic in nature. One molecule of ozone produced in the atmosphere is 1200-2000 times more effective in global warming than an additional CO₂ molecule (Lal 2007). It is a secondary pollutant photo chemically produced during the oxidation of methane (CH₄), carbon monoxide (CO) and non-methane volatile organic compounds (NMVOC) in the presence of nitrogen oxides (NOx = NO+NO₂) (Tang et al. 2009). The five main factors, which control the tropospheric ozone levels, are photo-chemical production, chemical destruction, longrange transport, surface dry deposition on vegetation, water and other materials and stratospheric-tropospheric exchanges (Alvim et al. 2010). Due to strong association between air pollutant levels and meteorological conditions, meteorological variability influences the air quality caused by changes in emissions of pollutants (Rao and Zurbenko 1994). Additionally chemical generation of pollutants and their deposition on the surface are strongly affected by meteorological variables (Tarasova et al. 2003). With this view, the influence of local meteorological conditions on tropospheric pollutants (tropospheric ozone, NO_2 and particulate matter, $\mathsf{PM}_{2.5}$ and PM₁₀) levels was studied at Dayalbagh, Agra from January 2011 to December 2011. In this study, an attempt has been made to address briefly the diurnal-seasonal variations in surface ozone,NO2 and particulate matter and the inter-relation of these pollutants and available meteorological parameters (temperature, wind speed, relative humidity).

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pollutants and available meteorological parameters (temperature, wind speed, relative humidity). Materials and Methods Sampling Site

Agra (27°10' N and 78°05' E) with an altitude of 169 m above sea level lies in the semi-arid zone of India and is positioned with the Thar Desert of Rajasthan to the West, central hot plains to the South, Gangetic plains to the East and cooler hilly regions to the North (Figure 1).



Figure 1: Map showing Geographical Location of the Sampling Site

Sampling was performed at Dayalbagh, a sub-urban area located in the north of Agra primarily a vegetated area with minimum industrial activity. The sampling site lies by the side of a road that carries mixed vehicular traffic, moderate (of the order of 1000 vehicles) during the day and minimal (of the order of 100 vehicles) at night. The campus lies about 2 km north of the National Highway-2 (NH-2), which has dense vehicular traffic (10⁶ vehicles) throughout the day and night (Singla et al. 2011).

Measurement of Ozone and NOx

Measurement of surface Ozone and NO2 was carried out at Dayalbagh from January 2011 to December 2011. O₃ concentrations were recorded using continuously operating O_3 analyzer (Thermo Fischer Model 49i) based on ultraviolet absorption photometry, resting upon absorption of radiation with the wavelength of 254 nm by ozone in the analyzed sample. Calibration of the system was done regularly with the help of a built-in ozone generator. Zero setting of the analyzer was also performed regularly by admitting zero air (free of ozone) into the analyzer. The minimum detection limit and precision of the analyzer are about 0.5 ppbv and 1 ppbv (2 %) respectively. NO_x concentrations were recorded with a NOx analyzer (Thermo Fischer model 49i) based on the chemiluminescence effect of NO₂ produced by the oxidation of NO by O3 molecules, which peaks at 630 nm radiation. In this analyzer, NO2 is measured by converting it into NO using the thermal conversion (heated molybdenum) method. The minimum detection limit and precision of NOx analyzer are 0.4 pbv and 1 ppbv (2%) respectively. Calibration for NO₂ was done by using permeation tube (KINTEK, TEXAS). The online O3 and NOx analyzers record concentrations on an averaging time of every 5 minutes. The 5 minute average concentrations were

further averaged to get one hour average O_3 and NOx concentration (Singla et al., 2011).

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Sampling of PM_{2.5} and PM₁₀

Concurrently $PM_{2.5}$ and PM_{10} were collected using a high volume sampler, Envirotech APM 550, installed on the roof of a building. The samples were collected on pre-desiccated and pre-weighed glass fiber filters (Whatman, EPM 2000). Air was drawn at a flow rate of 16.6 L min⁻¹ for 24 hr on an average for collection of both $PM_{2.5}$ and PM_{10} . Each filter was weighed at least thrice before and after sampling and the net mass was obtained by subtracting the presampling weights from the post-sampling weights. PM concentration was calculated by dividing the net mass by the volume of air sampled.

Recording of Meteorological Parameters

Ambient temperature, solar radiation, relative humidity, wind speed, and wind direction were recorded through an automatic weather monitoring system (WM251) mounted on the roof. It was programmed to collect data at 1 min interval and store them in memory to be downloaded to a computer. In summer months, the average temperature ranged from 40-45°C and average humidity varied from 25-50%. Summer was also associated with strong dust storms (wind speed varied from 10-16 kmh⁻¹). The average temperature varied from 22-28°C and relative humidity varied from 70-75% in winter months. During monsoon (July-September) relative humidity reached up to 80-85%.

Results and Discussion

Trends of Ozone and NO₂ Concentrations

Figure 2 displays the mean and 1o standard deviation of O₃ and NO₂ concentrations observed in each month during the study period (30/31 day average). The highest monthly mean **O**₃ concentrations were observed in summer season (April to June) followed by winter (December to February), post monsoon (October and November) and monsoon (July to September). The higher ozone concentrations varied from 96.5 to 129.3 µg m during summer season while the lower O_3 concentration varied from 11.9 to 31.9 µg m⁻³.





Figure 2: Mean Concentration levels of O₃ and NO₂ in different months

On the other hand during winter and post monsoon season, the average O_3 concentration (stretched from 46.04 to 96.79 $\mu g\ m^3$ in winter and 55.2 to 91.4 µg m⁻³ in post monsoon) was observed to be much lower than that observed in summer season. This might be due to peak surface temperature observed in summer reaching as high as 41 to 46°C and lowest seasonal total rainfall of 30.9 mm. The maximum O₃ concentration varied from 154.2 to 210.2 µg m⁻³ and minimum concentration varied from 8.78 to 21.6 µg m⁻³ in winter season. During post monsoon, the maximum concentration varied from 139.8 to 208.1 µg m⁻³ and minimum varied from 7.71 to 20.7 $\mu g m^{-3}$. Low concentrations of O₃ were observed in monsoon season (average concentration varied from 43.5 to 70.2 µg m⁻³). During monsoon season, the maximum levels varied from 108.7 to 146.3 µg m and minimum levels varied from 4.7 to 12.2 μ g m⁻³. These lower concentrations are attributed to scavenging of pollutants by rain. (Reddy et al., 2008). Similarly, higher concentrations of NO₂ were observed in summer, winter and post monsoon season with remarkably lower values in monsoon season. The average NO_2 concentration varied between 12.4 to17.8 µg m⁻³ in summer, 10.3 to 14.1 µg m⁻³ in winter, 7.6 to 11.5 μ g m⁻³ during post monsoon and 4.4 to 6.5 μ g m⁻³ in monsoon season. Relatively higher concentrations observed in winter might be due to temperature inversion phenomenon while lower concentrations in monsoon could be associated with rains and overcast sky (50-100%) conditions (www.wunderground.com).

Diurnal Variation

The variation of surface ozone, within a day, may be helpful in delineating the processes responsible for ozone formation at a particular location. A representative average diurnal pattern of O₃ and NO₂ is shown in Figure 3. As expected, time concentration profiles of O3 show a potential and direct link with NO2 variability. In general, O3 concentration starts increasing gradually after sunrise (0800 h), attains maximum concentration at noon hours and then gradually decreases and thus closely follows the diurnal cycle of solar radiation and surface temperature. However, NO₂ concentration shows a Asian Resonance

sharp increase and reaches a maximum at a time that coincides with the maximum automobile traffic in the morning and evening respectively and thus follows an inverse relationship with O₃. In the early morning hours before sunrise (0600 h), O₃ levels are observed to be very low and almost constant (16-22 µgm⁻³). With the onset of sunshine, as solar radiation increases, O₃ concentration gradually increases (39-79 µgm⁻³) probably due to its photochemical formation through photolysis of NO2 via following set of reactions.

 NO_2 + hv (λ <420 nm) \rightarrow NO + O $O+O_2 + M \rightarrow O_3 + M$

(2)

The concentration of O_3 that attains maximum values (125-129 µg m⁻³) at noontime (1200-1400 h) can be attributed to O3 production by photooxidation of precursor gases, like CO, CH₄, and other hydrocarbons in the presence of sufficient amount of NO/NOx. In the evening hours, a decrease in the O₃ values (24-29 µg m⁻³) is observed probably due to lower production of O₃ as well as its depletion due to the fast reaction with NO forming NO2 (O3 + $NO \rightarrow NO_2 + O_2$). This daytime increase in ozone concentration is a pronounced feature of an urban site (Lal et al. 2000, Reddy et al. 2008). Boundary layer processes and meteorological parameters also play a role in the O₃ variability. After sunrise, boundary layer height gradually increases from about 200-300 m to about 1500-2000 m during noon hours due to convective heating, and the layer becomes stratified (Reddy et al. 2008b). The photochemical production of O_3 is directly related to temperature and solar flux, its concentration increases with respect to both parameters.



Figure 3: Average Diurnal Variation of O3 and NO2

NO₂ mixing ratios are generally stronger during the morning and late evening hours. Variations in NO2 are caused by variations in boundary layer processes, chemistry, mixing anthropogenic emissions, and local surface wind patterns (Rao et al. 2002a Rao et al. 2002b). The emitted pollutants get trapped at the lower heights in the boundary layer during evening and continue to remain so until early morning due to formation of a nocturnal stable layer. However, pollutants are diluted during midday due to increased height of the boundary layer and extensive mixing. Vehicular emissions serve as the dominant source for NO₂ in this region. During the morning hours (0600-0800 h) the high values of NO2 (39-45 μ g m⁻³) have been observed and hence expected O₃ titration to become faster. This might be because of increase in heavy-duty diesel truck activities that serves as a significant source of NOx emissions. At evening traffic peak hours (2000 h), a secondary NO2 maximum (~39.3 µg m⁻³) is observed during which time O₃ exhibits a secondary minimum (~7.5-9 µg m ³). A diurnal pattern with maximum ozone concentration during daytime and minimum during nighttime is also observed during dry months at various locations in India like Ahmedabad, Delhi, Pune, Trivandrum, Anantapur and Gadanki (Purkait et al. 2009, Naja et al. 2003, Satsangi et al. 2004, Sahu and Lal 2006, Londhe et al. 2008, Debaje and Kakade 2009, Naja and Lal 2002).

PM_{2.5} and PM₁₀

The mass concentrations of both PM2.5 and PM₁₀ exhibited distinct seasonal variations similar to many other areas in the world (Zheng and Fang 2000). The mass concentrations displayed higher concentrations in winter than monsoon by a factor of 2. Figure 4 shows the monthly variation in PM_{2.5} and PM₁₀ mass concentrations. As observed from the graph, the mass concentration of PM2.5 varies between 58-84 µg m⁻³ in summer season (followed by a significant dip in concentrations in monsoon season where mass concentration varied from 23-32 µg m⁻³ Thereafter, the mass concentrations start rising rapidly in October and become highest in winter season with the concentration reaching between 105-217 μ g m⁻³. Likewise the mass concentration of PM₁₀ varied from 178-250 μ g m⁻³ in summer season followed by lower concentrations (42-79 µg m⁻³) in monsoon and markedly higher concentrations (110-324 µg m⁻³) in winter.

The seasonal distribution of particulate matter is proportionately governed by the combined effect of emission factors, dispersion conditions and chemical mechanisms. In Agra (India), summers are characterized by dry climate with intense dry storms and intermittent rainfall is observed in the monsoon while winters are usually associated with calm periods and stagnant atmospheric conditions. During winters, particulate emissions from automobile exhaust are higher due to low ambient temperature and traffic congestion due to poor visibility, fog, mist and haze. Also fuel combustion processes involving wood burning and coal combustion for cooking and space heating during winter add to the levels of particulate

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matter. The lower particulate concentrations in monsoon are attributed to their scavenging through precipitation and reduced soil re-suspension. The sources and transport pathways of particulate matter were identified using air mass backward trajectory analysis. These air mass back-trajectories were National calculated using the Oceanic and Atmospheric Administration (NOAA) Hybrid Single-Particle Lagrangian Integrated Trajectory (HYSPLIT) model (http://ready.arl.noaa.gov/HYSPLIT. php) which is based on the GDAS Meteorological Data (Draxler et al. 2003). The five-day back trajectories for different seasons (summer, monsoon and winter) were calculated at 14:00 h (local time) for 1000, 1500 and 2000 m above the ground level and have been represented in Figure 6.

In summer long-range transport of dust from dust storms originating from the Thar Desert contributes to higher concentrations of particulate matter. During monsoon season, the moisture loaded winds originating from Arabian Sea and Bay of Bengal causes heavy rainfall, thereby lowering particulate concentrations significantly. Short trajectories originating from shorter distances in winter season reflects the major contribution from local sources to higher concentrations of PM_{2.5} and PM₁₀.



Figure 6: The typical air mass five-day trajectories arriving at the sampling site in a) summer season pointing to the dust source from northwest Asian desert regions b) monsoon season indicating the origin of air masses from the Arabian Sea and Bay of Bengal and c) winter season indicating local origin of aerosols





Figure 4: Monthly Mean Concentration of PM2.5 and PM10

Relationship of Tropospheric Pollutants with Meteorological Parameters

Relationship of O3, NO2 and Particulate Matter with Temperature

The temperature variance observed during the study period was found enough to study the relationship between temperature and pollutants of interest. This was studied on a seasonal basis as shown in Table 1. The correlation between temperature and primary pollutant, NO2 was found to be insignificant in all seasons as compared to secondary pollutant, O₃. Surface ozone, an indicator of photo-oxidation processes exhibited positive trend with surface temperature. O₃ increases significantly with the increase in surface temperature and showed strong positive correlation with temperature in summer and a moderate correlation in winter (Table 1). Both PM_{2.5} and PM₁₀ showed an inverse correlation with temperature (Table 1). This negative correlation could be attributed to dispersion of particulate matter due to unstable conditions associated with high temperature as well as due to faster thermal degradation of some constituents of aerosols. In general, higher concentrations of PM_{2.5} and PM₁₀ were observed during days with temperature less than 20°C.

 Table 1: Correlation between Pollutants and Meteorological Variables

	Season	Temp	WS	RH
03	Summer	0.75*	0.52*	-0.65*
	Winter	0.57*	0.48*	-0.82*
	Monsoon	0.45	0.29	-0.26

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NO ₂	Summer	-0.43	0.55*	-0.49*		
	Winter	-0.55	0.67*	-0.31		
	Monsoon	-0.20	0.26	-0.25		
PM _{2.5}	Summer	-0.59*	-	0.74*		
			0.62*			
	Winter	-0.56*	-	0.81*		
			0.73*			
	Monsoon	0.12	-0.36	0.50*		
PM_{10}	Summer	-0.62*	-	0.68*		
			0.58*			
	Winter	-0.58*	-	0.76*		
			0.66*			
	Monsoon	0.12	0.10	0.52*		
orrelations are significant at 0.05 level (2 tailed)						

Relationship of O_3 , NO_2 and Particulate Matter with Wind Speed

The prevailing winds loaded with moisture or aerosol particles from a distant source play a major role and may affect the tropospheric pollutant concentration. The concentration level of O_3 and NO_2 were observed to increase with the transport of wind and hence showed positive correlation with wind speed. O_3 was found to be moderately correlated with wind speed and NO_2 showed considerable positive correlation (Table 1). On the other hand, the negative relationship between particulate matter and wind speed suggest that strong winds flush away the particulates from the study area and low winds allow particulate levels to rise. Concentrations of $PM_{2.5}$ greater than 100 µg m⁻³ were observed with wind speed less than 2 km h⁻¹. Similarly, PM_{10} concentrations were higher than 250 µg m⁻³ when wind speed was less than 2 km h⁻¹.

Relationship of O3, NO2 and Particulate Matter with Relative Humidity

 O_3 and NO_2 showed significant negative correlation with relative humidity (Table 1). This may be attributed to enhanced oxidation of hydrocarbons in the afternoon and thereby supports the ozone production (Ma and Van 2000). The reaction of NO_2 with OH is decreased at lower relative humidity and hence concentrations of NO_2 increase. On the other hand, $PM_{2.5}$ and PM_{10} showed moderately positive correlation with humidity.

O₃ Episode

Agra is a tropical site, experiences extreme climatic conditions in summer season and in the winter season as well. The site observes continuous 7-9 hr of sunshine during a day i.e. from 8:00 a.m. in the morning till 5:00 p.m. in the evening. Since solar radiation is the major factor responsible for the photochemical processes to occur, hence significantly high O_3 concentrations were observed during nine consecutive days from 25th May to 4th June 2011, reaching as high as 258 µg m³. Figure 5 plots the hourly variations of O_3 concentration, period of sunshine, temperature, relative humidity and wind speed for 9 consecutive days identified as episodic days in the summer of 2011.

Being in the tropical zone, the maximum temperature ranged from 40-45 °C. As observed from the figure 5, low temperatures (28-32 °C) are

observed in morning hours (6:00 - 8:00 a.m.) and thereafter starts rising and becomes maximum (45-46 °C) in noon hours (12:00 - 2:00 p.m.). Particularly, ground level ozone concentrations show positive correlation with surface temperature. Relative humidity varied from 25-60% during the episodic O3 days and shows significant negative relationship with O₃ concentration. O₃ destruction occurs whenever an excited oxygen atom $O(^{1}D)$ reacts with $H_{2}O$. Additionally, water vapour also enhances the removal of short-lived radicals such as HOx, which are important precursors for O₃ formation. Further, wind speed varied from 2.5 to 5.6 km h⁻¹ during night hours and 6.1 to 11.3 km h⁻¹ in the daytime hours i.e. from 10:00 a.m. to 4:00 p.m. Since low wind speed favors O₃ formation, therefore wind speed was observed to be moderately correlated with O₃ concentration.



Figure 5: Hourly variations in O3 concentration, period of sunshine, temperature, relative humidity and wind speed during high O3 days Conclusion

The diurnal and seasonal variation of O_3 displays noon-time (125-129 µg m⁻³) and summer season maxima (96.5 to 129.3 µg m⁻³) respectively while diurnal variation of NO₂ shows morning and evening maxima (39-45 µg m⁻³) corresponding to traffic rush hours. The mass concentration of PM_{2.5} (105-217 µg m⁻³) and PM₁₀ (110-324 µg m⁻³) was observed to be maximum in winter. Ozone was positively correlated with temperature while PM_{2.5} and PM₁₀ were negatively correlated with temperature. Solar radiation was also strongly related with O₃. Wind speed showed positive correlation with NO₂. O₃ and NO₂ showed significant negative trend with relative humidity respectively which may be due to enhanced oxidation of hydrocarbons in the afternoon. Acknowledgements

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